MINERALOGICAL STUDIES ON BONE APATITE AND THEIR IMPLICATIONS FOR RADIOCARBON DATING

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ABSTRACT. Infrared (IR) spectrophotometry and X-ray diffraction (XRD) were conducted on modern and fossil bone material from archaeological sites in the U S to determine post-mortem changes in bone apatite and to evaluate the effect of these changes on radiocarbon dating. IR absorption bands, XRD peak-broadening parameters, and XRD unit cell measurements indicated that during fossilization, bone apatite, a mineral similar to dahllite, was partially or completely recrystallized to francolite. Post-mortem changes involved then removal of some of the endogenous crystal carbonate both at surfaces and at internal OH-sites and introduction of exogenous carbonate into internal crystal PO₄ sites. Increased fluorine content accompanied carbonate substitution.

Both the carbon isotopic composition and the amount of exogenous carbonate introduced into the apatite structure will affect the radiocarbon dating of bone apatite. Special sample pre-treatment may remove most of the substituted carbonate in some cases. Simulated experiments are suggested for a better understanding of the nature and mechanism of carbonate substitution in bone apatite for the removal of the exogenously substituted carbon and the improvement of radiocarbon dates.

INTRODUCTION

Despite the abundance of bone material at archaeological sites, only a few radiocarbon dates have been obtained from them. Many workers regard bone dating with suspicion because of the inaccurate ages often obtained. The organic fraction in bone, namely collagen, does not exchange carbon with the bone post-mortem environment. Accordingly, inaccurate collagen dates are caused by physical contamination with other organic material, which has led many investigators to devise new techniques for separating pure collagen (Berger, Horney, and Libby, 1964; Krueger, 1965; Sellstedt and others, 1966; Berger and others, 1971; Longin, 1971; and Hassan, 1976). In many cases, however, the amount of collagen separated is insufficient for dating. Haynes (1968) used the bone mineral, apatite, as an alternative.

Many investigators are skeptical of radiocarbon dates obtained from bone apatite because of possible post-mortem exchange of carbon with the surrounding medium (Tamers & Pearson, 1965; Sellstedt and others, 1966). One of us (CVH), in collaboration with Austin Long, University of Arizona (unpub), demonstrated this exchange by radiocarbon analysis of about 2 my old bone from the Curtis Ranch locality in Arizona and obtained apparent ages of $21,000 \pm 1000$ BP (A-1028) on apatite CO₂ and $16,180 \pm 300$ BP (A-1027) on CO₂ from secondary carbonate. The mechanism and magnitude of such exchange had not been studied before. In the present study, modern and fossil bone apatites were subjected to X-ray and infrared investigation to trace the changes which might have

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TABLE 1
Frequency for IR absorption bands (in cm⁻¹) for the carbonate in dahllite, francolite, aragonite, calcite and bone apatite

			•	
Samples	$ u_1$	ν_2	ν_3	ν_4
Dahllite				
(Carlström, in press)		873	1416	
- '			1455	
Francolite				
(Carlström, in press)		866	1425	
• ,			1455	
Aragonite				
(LeGeros and others, in press)	1080	842	1445	700
• /		852	1485	710
Calcite				
(LeGeros and others, in press)		872	1420	710
Protein-free rat tibia, untreated				
(Termine and Lundy, 1973)		868	1417	
(,,)		877*	1441*	
			1476	
Protein-free rat tibia, air, 24 hr				
(Termine and Lundy, 1973)		870	1413	
,, ,,		876	1457	
			1485*	

^{*} Shoulder

occurred during fossilization and their possible effects on radiocarbon dating. The fossil bone specimens are from Folsom site, New Mexico; Murray Springs Clovis Level, Arizona; Blackwater Draw Clovis Level, New Mexico; Domebo Clovis Level, Oklahoma; and Boney Springs and Trolinger Bog, Missouri. They belong to *B bison antiquus, Mammuthus columbi, Mammuthus, imperator,* and *Mammut americanum*. The modern equivalents used in the study consist of buffalo (bison), cow, and elephant (*E indictus*)¹. The elephant bones belong to two individuals; one died in 1967 and was left in open air while the other died in 1964 and was buried for 6 years prior to re-excavation.

INFRARED SPECTROSCOPY

Infrared (IR) absorption spectra from bone are usually composites of both protein and mineral absorption bands. The spectra of bone apatite are characterized by vibration bands produced by H_2O , OH, PO_4 , and CO_3 groups. The characteristic vibrational bands of the CO_3 group in spectra of bone apatite are different from those observed in spectra of simple carbonate (table 1). The CO_3 IR bands of fresh bone apatite (table 1) are more complex than those of dahllite (carbonate hydroxyapatite with less than 1 percent F^-) and francolite (carbonate fluorapatite with more than 1 percent F^-).

Generally CO₃ exists in bone apatite at three different positions: (1) absorbed on the crystal surface, (2) at OH-, and (3) at PO₄-sites

 $^{^{\}mathtt{1}}$ The elephant bones were obtained from Charles McNulty, University of Texas at Arlington.

Sample description	Ca%	PO ₄ %	$CO_3\%$	F%
	36.18	53.96		.074
Fresh cow	36.06	52.93	5.75	.035
Fresh buffalo	35.04	49.13		.076
Modern elephant Modern buried elephant	34.94	48.93	6.50	.055
Folsom bison	39.08	50.24	5.80	1.404
Murray Springs mammoth	37.46	49.28		1.467
Blackwater Draw mammoth	37.59	49.89	6.56	1.830
Blackwater Draw mammoth III	37.16	45.24		3.958
Domebo mammoth	36.67	48.44	9.78	1.659
Boney Springs mastodon	38.70	49.35		1.561
Trolinger bog mastodon	38.88	50.50	5.52	1.781

Table 2
Chemical analysis of fresh, modern, and fossil bone apatite*

(LeGeros and others, 1967; LeGeros and others, 1969; Carlström, in press; and Elliott, in press). About one-third to two-thirds of the total carbonate occurs at the surface (Neuman & Mulryan, 1967; Termine & Eanes, 1972; Termine & Lundy, 1973). The remainder is distributed between PO₄- and OH-sites with the vast majority in PO₄-sites.

The spectra of bone apatite were made with a Perkin-Elmer Model 621 Spectrophotometer, purged with dry air from pellets of 300mg kBr and 2mg protein- and calcite-free samples.2 Some samples were treated with triammonium citrate and acetic acid to test the effect of these chemicals on carbonate removal. Machine adjustment was as follows: slit program 1x, gain 4.5, attenuator 11.00, scan time 10×4.5, suppression 6, and source 0.8 amp. The infrared splitting function SP, which is proportional to percentage crystallinity as determined by XRD, was calculated according to Termine and Posner (1966). The carbonate percentage was calculated from the ratio A₁/A₂, the ratio of the areas under the peaks of the carbonate ion asymmetric stretching frequency (1550-1350 cm⁻¹) and the phosphate ion antisymmetric bending frequency (800-450 cm⁻¹), and a standard whose CO₃ content is known (Greenfield & Eanes, 1972). The standard used in the present study is the modern buried elephant femur in which the CO₃ content was previously determined by wet chemistry (Hassan, 1976). Some of the chemical analysis on the bones investigated appear in table 2.

The spectra (fig 1) and frequencies (table 3) for IR linear absorbance for bone indicate that during fossilization bone apatite had experienced some changes. The absorption at 1495 cm⁻¹, attributed to surface and/or amorphous carbonate, is absent or very weak in the fossil bone. The intensities of IR bands at 1542, 1445, and 880 cm⁻¹, which stand for CO₃-OH substitution, decrease while an increase occurs at the 1418 and 873 cm⁻¹ bands, which stand for CO₃-PO₄ substitution, with fossilization. Despite these changes, the fossil bone apatites were still close to the dahllite-type of apatite with the exception of the bone of the Blackwater Draw Mammoth III, a francolite-type. It is noteworthy that this sample

^{*} From Hassan (1976)

² See Hassan (1976) for sample pretreatment.

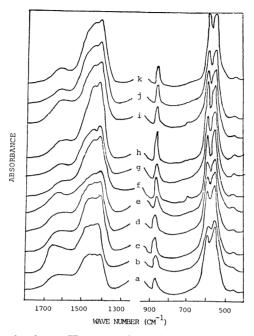


Fig. 1. Linear absorbance IR spectra from the bone of (A) fresh cow, (B) fresh buffalo, (C) modern elephant, (D) modern buried elephant, (E) Folsom bison, (F) Murray Springs mammoth, (G) Blackwater Draw mammoth, (H) Blackwater Draw mammoth III, and (K) Trolinger bog mastodon.

has an abnormally high F⁻ concentration (3.96 percent) whereas the rest of the fossil bones have less than 1.9 percent and the fresh and modern bones contain less than 0.1 percent flouride (table 2). The fact that bone mineral dahllite may be converted into francolite because of fossilization, has been reported previously by Stevenson (1966).

The relative carbonate content of fossil bone apatite decreases sharply with triammonium citrate and acetic acid treatments (table 4). Prior to these chemical treatments, calcite was removed mechanically (Hassan, 1976). Any calcite remaining was below the detection limit of the IR, about 2 percent. The decrease in CO₃ after the above treatments may well result from the weakening effect of the post-burial-substituted carbonate on bonds within the apatite structure. This is in agreement with the explanation of LeGeros and others (1967) for the observed increase in rates of dissolution and solubility of apatites with carbonate substitution.

X-RAY STUDIES

X-ray diffraction (XRD) broadening analysis has been extensively employed in the study of biological apatite (Carlström, 1955; Carlström and Glas, 1959; Posner and others, 1963, 1965; Harper and Posner, 1966; Termine and Posner, 1967; LeGeros and others, 1967; Posner, 1969;

TABLE 3 Frequency for IR absorption bands (in cm⁻¹) for the carbonate in modern and fossil bone apatite

	No treatm		Triammo citra		Acetic acid		
Sample description	ν_2	ν ₃	ν_2	ν_3	ν_2	<i>v</i> ₃	
Fresh cow	873 880*	1417 1445 1470 1495* 1512*	_				
Fresh buffalo	873 880*	1418 1445 1470 1495* 1512*		_ _ _ _			
Modern elephant	873 880*	1418 1445 1470 1500*		_ 	_	_ _ _	
Modern buried elephant	873 880*	1418 1445 1470 1495* 1510*	873 880*	1418 1445 1470 1495* 1510*			
Folsom bison	868 878*	1417 1450 1495*			_	_	
Murray Springs mammoth	873 880**	1423 1453 1495***	873 880***	1423 1453 1495*	873 880***	1423 1453 1495*	
Blackwater Draw mammoth	873 880**	1423 1453 1495***	873 880**	1423 1453 1495*	_		
Blackwater Draw mammoth III	866	1427 1455	866	$\frac{1429}{1455}$	-	_	
Domebo mammoth	873 880***	1420 1453 1495*	873 880*	1420 1453 1495*	873 880***	1420 1453 1495*	
Boney Springs mastodon	873 880***	1420 1453 1495*	_		_		
Trolinger bog mastodon	873 880***	1422 1453 1495*			_	_ _ _	

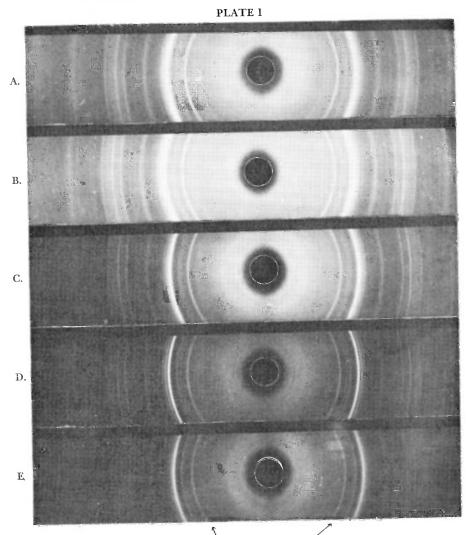
^{*} Shoulder ** Very weak shoulder *** Weak shoulder

Lundy and Eanes, 1973; and Russell, Termine, and Arcoli, 1973). Crystal-lite size can be deduced from the broadened X-ray diffraction peaks using Scherrer's equation (Klug and Alexander, 1954). However, the figures for bone apatite crystallite dimensions and its crystallographic orientation are controversial (Carlström and Glas, 1959; Posner and others, 1963; and Höhling, Themann, and Vahl, 1966) which suggests that the validity of accurate crystallite size determination for bone apatite has not been settled. However, it is generally agreed that the crystals are extremely minute with the smallest dimension less than 50 Å and the largest dimension much less than 1000 Å, paralleling the c-axis (Eanes, 1973).

Variability in unit cell dimensions of biological apatites resulting from chemical substitution and the mechanism of such substitution has also been a matter of controversy. Among the substitutions which may occur in the apatite structure is that of carbonate. Carbonate substitution was investigated by McConnell (1938, 1952a and b, 1959, 1962, 1965), Borneman-Starinkevich and Belov (1953), Trautz (1960), McClellan and Lehr (1969), Elliott (1965, in press), LeGeros (1965), LeGeros and others (1967), LeGeros and others (1968, in press), Simpson (1972), and Carlström (in press). Generally agreed is that in francolite, one carbonate and one fluorine substitute together for one phosphate in the apatite structure resulting in a nonlinear decrease in the a-axis. Experimental studies on bone powder and on synthetic apatite by Neuman and Mulryan (1949, 1967), Neuman and others (1949), Neuman (1950), Neuman and Weikel (1954), and Simpson (1967) indicate that surface exchange, ionic and isotopic, as well as recrystallization, may also occur in bone apatites.

Table 4
Splitting functions (SP), the ratio A_1/A_2 , and carbonate percentage calculated from A_1/A_2 using the wet chemical analysis for the modern buried elephant as a standard

	No treatment			Tı	iammoni citrate	um	A	Acetic acid		
Sample description	SP	A_1/A_2	$CO_3\%$	SP	A_1/A_2	CO ₃ %	SP	A_1/A_2	CO ₃ %	
Fresh cow	0.042	0.880	4.3				,			
Fresh buffalo	0.033	1.350	6.6	_					_	
Modern elephant	0.036	1.314	6.4				_	_		
Modern buried	0,000	1.011	0.1				_			
elephant	0.023	1.327	6.5	0.023	1.296	6.3				
Folsom bison	0.047	1.345	6.6	0.040	1.450	0.5				
Murray Springs		-10-20	0.0					_		
mammoth	0.086	1.568	7.7	0.057	1.064	5.2	0.076	0.500		
Blackwater Draw		41000		0.037	1.004	5.4	0.076	0.708	3.5	
mammoth	0.059	1.327	6.5	0.047	1.067	5.2				
Blackwater Draw		-1041	0.0	0.017	1.007	5.4		_		
mammoth III	0.075	1.947	9.5	0.054	1.643	8.0				
Domebo mammoth	0.057	1.515	7.4	0.038	1.048	5.1	0.040	0.050	-	
Boney Springs		11010	7.1	0.036	1.040	5.1	0.049	0.870	4.3	
mastondon	0.039	1.452	7.1							
Trolinger bog					_	_	_		_	
mastodon	0.032	1.525	7.5							



 $4\theta = 51.85$

Exact positives of X-ray films of (A) fresh cow bone, (B) modern buried elephant bone, (C) Folsom bison bone, (D) Blackwater Draw Mammoth III bone, and (E) Domebo mammoth bone.

XRD peak broadening analyses were obtained from protein-free finely hand-ground samples³ adapting the method and experimental conditions of Russell, Termine, and Arcoli (1973). The apatite c-axis (002) and a-axis (310) reflections were step-scanned between 23.60-27.60° 2θ and 36.00-44.00° 2θ , respectively. Russell, Termine, and Arcoli (1973) ob-

³ See Hassan (1976) for sample pre-treatment.

 $\label{eq:Table 5} \text{X-ray diffraction broadening and intensity parameters}$

Sample description	B ₀₀₂ 29	B ₃₁₀ 20	1/2°28 B310°28 B002 8 B31028 EB1°28	B310 ² 9	ΣΒ ^{1°} 2θ	1 ₀₀₂ (area)	I ₃₁₀ (area)	Brass Count	I/Brass 002	1/Brass 310	l xin	DA 002	DA.310
Fresh cow	0.41	1.29	0.54	1.60	2.14	21,251	29,372	5,851	3.632	5.020	8.652	371	91
Fresh buffalo	0.39	1.31	0.49	1.66	2.15	20,372	31,663	5,844	3.486	5.418	8.904	408	89
Modern elephant	07.0	1.28	0.52	1.56	2.08	21,575	29,739	5,821	3.706	5.109	8.815	388	92
Modern buried elephant	0.39	1.23	67.0	1.53	2.02	21,705	30,390	5,874	3.695	5.174	8.869	807	26
Folsom bison	0.35	1.02	0.45	1.33	1.78	23,750	34,306	5,776	4.112	5.939	10.021	510	128
Murray Springs mammoth	0.32	0.95	0.39	1.09	1.48	25,144	34,615	5,867	4.286	5.890	10.176	628	143
Blackwater Draw mammoth	0.32	96.0	0.40	1.20	1.60	30,884	35,419	5,831	5.297	6.074	11.371	628	141
Blackwater Draw mammoth III	0.29	0.49	0.36	0.97	1.33	28,006	39,081	5,818	4.814	6.717	11.531	816	650
Boney Springs mastodon	0.35	1.11	0.43	1.36	1.79	29,932	34,787	5,794	5.166	6.004	11.170	510	113
Trolinger bog mastodon	0.36	1.20	0.45	1.48	1.93	31,473	36,354	5,921	5.316	6.140	11.456	480	101
Fluor-apatite ^a Standard	0.19	0.36				1			1	I			

a obtained from Mr. B. O. Fowler at NIH.

tained a precision in half-width (B 1/2) and integral breadth (B^I = peak area/peak height) values within ± 1.5 percent. Debye-Scherrer, 114.6mm diameter powder cameras were used for study of unit cell dimensions. With Ni-filtered Cu radiation (CuK₂ = 1.542 Å), XRD patterns were recorded at 35 kv and 20 ma. The exposure time was 2 hrs. The 002 and 300 X-ray diffraction lines were used to calculate c Å and a Å dimensions, respectively.

The crystallite dimensions, D Å₀₀₂ and D Å₃₁₀, of modern bone apatite obtained in the present study (table 5) have their small dimension parallel to the a-axis, in agreement with Carlström and Glas (1959). The crystallite size of bone apatite and/or the percent crystallinity increases with fossilization. The XRD lines (pl. 1) are sharper in fossil than modern bones which most probably results from crystallite enlargement. The calculated unit cell dimensions (table 6), even though not very accurate, show a distinction between modern and fossil bone apatite. The slight decrease in the a-axis with fossilization is consistent with the observed increase in fluorine and carbonate content detected with fossilization (tables 2 and 4).

CONCLUSIONS AND PROBLEMS

Fossil bone apatite undergoes a number of post-mortem changes which distinguishes it from its biological predecessor. Among those changes critical for radiocarbon dating is carbonate substitution. The present study reveals a definite change in the mode and amount of carbonate contained within bone apatite upon fossilization which suggests post-mortem exchange with exogenous carbon. Whether the exchange is ionic and/or isotopic is still unclear. Chemical treatments with triammonium citrate and/or acetic acid remove some of the carbonate in the apatite structure. Carbon isotopic studies (Hassan, 1976) indicate that the carbonates thus removed are contaminants. Hydrochloric acid leaching of the finely ground samples (fractional hydrolysis technique) removes more of the contaminated carbonate which then may improve the radiocarbon dates (Hassan, 1976).

Table 6
Unit cell measurements

Sample description	θ_{002}	θ_{300}	d_{002}	d_{300}	c Å	a Å
Fresh cow femur	12.9500	16.5750	3.4371	2.7001	6.87	9.35
Modern buried elephant femur	13.0125	16.6000	3.4215	2.6961	6.84	9.34
Folsom bison bone	13.0000	16.6125	3.4241	2.6945	6.85	9.33
Murray Springs mammoth bone	13.0250	16.6250	3.4176	2.6922	6.84	9.33
Blackwater Draw mammoth III bone	12.9625	16.6500	3.4345	2.6882	6.87	9.31
Domebo mammoth bone	12.9625	16.6500	3.4345	2.6882	6.87	9.31
Boney Springs mastodon bone	13.0250	16.6750	3.4176	2.6843	6.84	9.30
Trolinger bog mastodon bone	12.9825	16.6250	3.4393	2.6922	6.86	9.33
Francolite*	-				6.881	9.356
Dahllite*					6.890	9.391
Bone**	_				6.88 ± 0.01	9.38 ± 0.01

^{*} Trautz, 1955

^{**} Posner and others, 1963

In order to evaluate the magnitude and mechanism of post-mortem exchange of carbon in bone apatite, simulated lab experiments are essential. Carbonate substitution in both very old and fresh bone powder using either active or dead carbon should be investigated. Data derived from such studies test the validity of the above mentioned fractional hydrolysis and demonstrate the limitations of its applicability to this problem.

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