YOUNG, OLD, AND WEATHERED CARBON—PART 2: USING RADIOCARBON AND STABLE ISOTOPES TO IDENTIFY TERRESTRIAL CARBON SUPPORT OF THE FOOD WEB IN AN ALKALINE, HUMIC LAKE

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ABSTRACT. Carbon (C) and nitrogen (N) stable isotope analysis (SIA) has been used to identify the terrestrial subsidy of freshwater food webs. However, SIA fails to differentiate between the contributions of old and recently fixed terrestrial C and consequently cannot fully determine the source, age, and biochemical quality of terrestrial carbon. Natural abundance radiocarbon (Δ^{14} C) was used to examine the age and origin of carbon in Lower Lough Erne, Northern Ireland. ¹⁴C and stable isotope values were obtained from invertebrate, algae, and fish samples, and the results indicate that terrestrial organic C is evident at all trophic levels. High winter δ^{15} N values in calanoid zooplankton (δ^{15} N = 24‰) relative to phytoplankton and particulate organic matter (δ^{15} N = 6‰ and 12‰, respectively) may reflect several microbial trophic levels between terrestrial C and calanoid invertebrates. Winter and summer calanoid Δ^{14} C values show a seasonal switch between autochthonous and terrestrial carbon sources. Fish Δ^{14} C values indicate terrestrial support at the highest trophic levels in littoral and pelagic food webs. ¹⁴C therefore is useful in attributing the source of carbon in freshwater in addition to tracing the pathway of terrestrial carbon through the food web.

INTRODUCTION

The investigation of trophic levels in a food web is complex yet fundamental to understand the dynamics of a freshwater system. Identification of carbon sources can be problematic due to large variation of stable carbon isotope values in freshwater systems, and becomes even more difficult when there is a significant terrestrial subsidy of the food web as terrestrial and freshwater isotope values often overlap. Carbon inputs to lakes can represent a significant component of the diet of lake heterotrophs (organisms unable to synthesize their own organic carbon).

Organic carbon utilized in aquatic food webs is derived from both autochthonous (derived from primary production within the system) and allochthonous (terrestrial) carbon imported from the catchment. Studies showing the influence of terrestrial carbon on food web dynamics have largely dealt with non-eutrophic lakes, which are predominantly supported by terrestrial carbon (Ågren et al. 2008; Gratton et al. 2008; Seifert and Scheu 2012; Lapierre et al. 2013; Wilkinson et al. 2013a,b). However, lakes with higher trophic status may also be supported by catchment inputs; the terrestrial support of pelagic food webs in such lakes is difficult to ascertain and can be underestimated.

Seasonal changes in productivity affect the delivery of nutrients (Grey et al. 2001; Kritzberg et al. 2006b; Neff et al. 2006; Rautio et al. 2011). Terrestrial carbon can supplant autochthonous carbon in systems with low algal abundance. The availability of this carbon to lake heterotrophs (organisms that cannot fix carbon but utilize organic carbon for growth) is related to its age, which in turn is related to its biochemical composition and nutrient quality (biochemical quality). Bacteria have been shown to metabolize both autochthonous carbon and labile and decayed terrestrial carbon (Guillemette et al. 2013) but preference is linked to nutrient content, which decreases over time, leading to preferential selection of autochthonous and labile terrestrial carbon sources (Taipale et al. 2007, 2008; McCallister and del Giorgio 2008; de Kluijver et al. 2012; Zigah et al. 2012).

Freshwater Carbon Sources

Carbon derived from primary production is bio-energetically rich compared to terrestrial input to lakes, particularly previously sequestered detrital carbon that has undergone diagenesis in soil or

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peat. As carbon sources fluctuate seasonally, heterotrophic organisms can switch between autochthonous or allochthonous sources to exploit high-quality carbon sources (Maguire and Grey 2006; Rautio and Vincent 2007; Taipale et al. 2008; Rautio et al. 2011).

Terrestrial organic matter has been shown to enter aquatic food webs in a number of ways. Terrestrial dissolved organic carbon (DOC) can be utilized by mixotrophic (utilizing a mixture of autotrophic and heterotrophic modes) and heterotrophic bacteria (Tittel et al. 2009; Cole et al. 2011; Rottberger et al. 2013). Particulate organic matter (POM) of terrestrial origin often forms important foci for bacterial production. Subsequent consumption of bacteria can also be an important pathway into food webs (Kritzberg et al. 2006a; Cole et al. 2011).

Stable Isotope Analysis of Food Webs

Food web utilization of terrestrial organic subsidies in lakes has been demonstrated for many widely distributed global locations. For the most part, these involve the use of natural abundance or artificially enriched stable isotopes of carbon and nitrogen (Perga and Gerdeaux 2004; Carpenter et al. 2005; Kritzberg et al. 2006a; Marty and Planas 2008; Taipale et al. 2008).

Detecting the source of carbon can be problematic, however, when using stable isotope analysis. Kinetic stable isotope fractionation leads to differences in pelagic (open water) and littoral (near shore) areas of the lake, and between organisms and their dietary sources (Fry 1991). It is difficult to physically isolate phytoplankton samples and to accurately measure the baseline isotopic signature required in models detecting the carbon source utilized in food web studies (Hamilton et al. 2005; Cole et al. 2011; Smyntek et al. 2012). Isotope signatures of terrestrial, littoral, and pelagic sources can overlap considerably, preventing source discrimination. In addition, mobile predators can integrate carbon from a number of sources, resulting in a mixed dietary pattern that can be difficult to explain (Layman et al. 2012; Phillips et al. 2014). Allochthony in higher trophic levels may therefore not be detected. While the terrestrial subsidy of food webs can sometimes be established, the age (and therefore biochemical/nutrient quality) of carbon cannot be determined using stable isotopes alone.

Radiocarbon in the Food Web

Radiocarbon is useful for examining freshwater carbon dynamics. Δ^{14} C (see Methods for derivation) of freshwater carbon is governed by its source and is corrected for fractionation, removing the problem of kinetic fractionation experienced in stable isotope analyses. "Modern/recent" (atmospheric, modern photosynthetic terrestrial plant production) ¹⁴C will be equivalent to that of the modern atmosphere. "Labile" terrestrial carbon can be derived from terrestrial material that has undergone minor decomposition. Complex soil/peat dynamics can mean that although the carbon is chemically labile, it may have a lower than modern ¹⁴C content (Trumbore 2009). Therefore, carbon that is ¹⁴C-depleted relative to modern values may still be labile, despite its age, and available for microbial catabolization.

Due to soil carbon cycling and turnover, carbon can take decades to be released from subsurface soil (Trumbore et al. 1989; Polsenaere et al. 2013). As such, carbon is released that contains "bomb" ¹⁴C, enriched relative to the modern atmosphere due to nuclear testing in the 1950s and 1960s, resulting in positive Δ^{14} C values (Reimer et al. 2004).

"Subfossil" carbon originates from stored terrestrial soil/peat carbon that can be substantially ¹⁴C-depleted depending on its age and mineralogy, which may be significantly degraded (Neff et al. 2006; Guo et al. 2007; Raymond et al. 2007, 2013; Caraco et al. 2010; Butman et al. 2012).

The Δ^{14} C value of dissolved inorganic carbon (DIC) (and autotrophic producers utilizing DIC) may be partially derived from the weathering of carbonaceous bedrock (via atmospheric CO₂ or soil carbon) with no ¹⁴C content ("¹⁴C dead"). This leads to depleted ¹⁴C values and ¹⁴C ages older than the contemporary atmosphere. This difference is known as the freshwater reservoir offset (FRO) (Deevey et al. 1954; Broecker and Orr 1958; Broecker and Walton 1959; Ascough et al. 2007, 2010; Keaveney and Reimer 2012).

¹⁴C and stable isotope measurements of autotrophs and their consumers can be therefore be used to determine the age and contributions of autochthonous and terrestrial carbon to the food web. A study was carried out using ¹⁴C and stable isotope analysis to identify the carbon sources utilized in a complex lake to determine the age and influence of terrestrial carbon subsidies of the food web.

STUDY SITE AND METHODS

Study Site

Carbon dynamics in Lower Lough Erne (Figure 1) are complicated by seasonal changes in terrestrial subsidies and productivity (Keaveney et al. 2015 [this issue]). The lake was chosen as a suitable study site to build on the substantial amount of research carried out in the lake (Maguire and Gibson 2010). A ¹⁴C study (using fish and water samples) was carried out in the lake that showed an FRO of ~500 ¹⁴C yr was present due to the alkalinity of the lake (Keaveney and Reimer 2012). While the FRO was relatively large, it was lower than would have been predicted from the hardness of the lake water, most likely due to the contribution of younger allochthonous organic carbon in the lacustrine food web.



Figure 1 Lower Lough Erne, Fermanagh, Northern Ireland. Lake and river sites that were sampled are marked on the map. LE1, LE3, and LEP are pelagic sites; LEA1 to LEC are littoral sites.

Lough Erne in northwest Ireland is a system composed of two interconnected lakes. The limnology of the system has been well studied since the 1970s (Gibson et al. 1980, 1993, 2003, 2005; Zhou et al. 2000; Maguire and Gibson 2010). It is a relatively large, humic (DOC ~12 mg L⁻¹) and alkaline lake (alkalinity ≈ 1.5 meq L⁻¹). The lake is eutrophic on the basis of total phosphorous (TP)

(>50 μ g L⁻¹ since 1975; Battarbee 1986; Foy et al. 1993; Maguire and Gibson 2010), but light limitation due to high DOC loading results in Secchi depths of 1–3 m.

Chlorophyll *a* concentrations were reduced due to the presence of zebra mussels (*Dreissena morpha*) (Foy et al. 1993; Maguire and Gibson 2010) since the late 1990s (Rosell et al. 1998). The expanding zebra mussel population with associated decreases in phytoplankton abundance have been shown by stable isotope analysis to shift Lower Lough Erne zooplankton production towards greater allochthony (Maguire and Grey 2006). There is no indication in these studies whether this terrestrial carbon is derived from modern or detrital material.

Field Sampling

Field sampling was conducted in February, March, May, August, and October of 2011 and details of methodology are described in the <u>online supplemental information</u> [also Keaveney et al. 2015 (this issue)]. The lake did not stratify during the study. DIC, DOC, POM, phytoplankton, and zoo-plankton were isolated from water samples collected at sites LE1, LE3, and LEP (Figure 1, Table 1).

of fish net placement.		
Site ID	Coordinates	
River Bannagh	54°32′02″N, 7°46′55″W	
LE1	54°29′07″N, 7°50′37″W	
LE3	54°23′51″N, 7°40′12″W	
LEP	54°21′30″N, 7°39′40″W	
LEA1 and 2	54°23′09″N, 7°40′08″W	
LEB	54°28′20″N, 7°51′51″W	
LEC	54°24′38″N, 7°41′50″W	
	Irish grid reference	Depth
K30	025 607	1 to 4 m
Pollan survey (PS)	110 610 to 130 600	20 to 30 m
	060 590 to 130 580	

Table 1 Location data for sampling sites including depths of fish net placement.

Littoral invertebrates were sampled once in July 2011 at three sites (Figure 1), using hand searches and 300-µm dip nets. Biofilm was collected from macrophyte stems and the upper surface of rocks using a toothbrush with matter rinsed into sample containers using Milli-Q[®] water.

Fish samples were obtained from routine surveys conducted for population and biomass studies at a number of spatially distinct sites carried out by the Agricultural and Food Biosciences Institute. Fish were caught using gill nets in 2010 using the survey methodology described by Rosell and Gibson (2000) and were stored frozen prior to sampling for isotope samples.

Laboratory Sample Preparation

Samples were pretreated and combusted for ¹⁴C and stable isotope analysis as described in the companion paper in this issue (Keaveney et al. 2015). The opercular bone obtained from fish samples was demineralized using 4% HCl for approximately 24 hr until CO_2 evolution was no longer evident. Paired opercular bones were homogenized using a Retsch MM200 Mixer Mill for stable isotope analysis; lipid removal was deemed unnecessary based on previous work in the lake (Keaveney and Reimer 2012).

Isotopic Measurements

The ¹⁴C/¹²C ratio of the sample relative to an international standard (F¹⁴C) and its associated uncertainty were calculated according to Reimer et al. (2004) and van der Plicht and Hogg (2006) and incorporated a fractionation correction (Stuiver and Polach 1977) based on δ^{13} C measured by accelerator mass spectrometry (AMS). Δ^{14} C is defined in Stuiver and Polach (1977) as the relative difference between the absolute international standard (base year 1950) and sample ¹⁴C/¹²C ratio corrected for age and δ^{13} C fractionation. The Δ^{14} C age correction accounts for decay that took place between collection/death and the time of measurement, so that two measurements of the same sample taken years apart will produce the same Δ^{14} C result, which is calculated using the equation Δ^{14} C = [F¹⁴C * e^{λ (1950-x)} - 1] * 1000, where *x* is the year of collection or growth (2010 for the fish samples, 2011 for all other samples).

The AMS δ^{13} C measurements (not given) include any fractionation that may occur during sample preparation and AMS analysis and therefore are only used for fractionation correction. All other stable isotope measurements were made using a Thermo Scientific EA-IRMS. Samples were corrected using a one-point calibration (R041 L-alanine). Other standards were measured at intervals (USGS 41 and leucine) and results plotted to make a two-point calibration. The effect of applying one-point and two-point correction to samples was compared and the difference was negligible. The precision of the EA-IRMS is <0.1‰ for δ^{13} C and <0.15‰ for δ^{15} N. Isotope values were expressed relative to the known standards ($\delta^{13}C_{VPDB}$ and $\delta^{15}N_{ARB}$).

Statistical Analyses

To determine the percentage contribution of sources to the diet of Lower Lough Erne fish, mixing models were generated using SIAR v 4.2 (Parnell et al. 2008), a Bayesian isotope mixing model in R (v 2.12.1). The model utilizes mean stable isotope values and associated standard deviations from dietary sources to generate a mixing model that incorporates these uncertainties associated with multiple sources.

Trophic enrichment factors (TEF) were $\Delta^{13}C = 1.63 \pm 0.5\%$ and $\Delta^{15}N = 2.9 \pm 0.32\%$ (Caut et al. 2009). SIAR models have been criticized for sensitivity to variable trophic enrichment factors, which in this study were obtained from the literature (Caut et al. 2009; Bond and Diamond 2011; Phillips et al. 2014). Changes made to the TEF when generating SIAR models with data from this study had little influence on the results of the model. However, large differences in stable isotope signatures can make the model less sensitive to trophic enrichment factors (Jensen et al. 2012). As such, the model was thought unreliable for ¹⁴C data and was only utilized and assumed to be robust for stable isotope data.

Models were generated from measurements of the largest fish groups (perch: *Perca fluviatilis*; pollan: *Coregonus autumnalis*; and roach: *Rutilis rutilis*) and source prey collected from Lower Lough Erne in 2011. Data utilized in models may be found in the <u>supplemental information</u>. All other statistics (regression, ANOVA) were performed in R using the HMisc (version 3.14-0, Harrell et al. 2007) and ggplot (version 0.9.3.1, Ginestet 2011) packages.

RESULTS

Lower Lough Erne Fish

Stable isotope analyses of Lower Lough Erne fish show distinct differences between species (Figure 2). Samples from deep water surveys (20–30 m, Table 1)—pollan (*Coregonus autumnalis*, n = 11) and trout (*Salmo trutta*, n = 3)—were ¹³C-depleted ($\delta^{13}C < -28\%$). Elevated $\delta^{15}N$ (>17‰)

suggests the fish are more predatory than others measured in the study. Perch (*Perca fluviatilis*, n = 16) collected in the same location had similarly elevated δ^{15} N, yet perch were also more ¹³C-enriched than pollan and trout. The δ^{13} C values of perch collected in a shallower, more littoral sample (depth <4 m) vary from -29.0 to -21.5%; δ^{15} N values also differ (13.1 to 18.6%). Perch stable isotope results indicate a more mixed diet than pollan or trout with smaller fish feeding in the littoral area. Roach (*Rutilis rutilis*, n = 26) isotopes also fluctuate showing a mixed diet with a littoral influence.



Figure 2 Stable carbon and nitrogen isotope results from Lower Lough Erne fish species collected in 2010

Based on the outputs generated in the mixing model, sources (labeled in Figure 3, detailed values in <u>supplemental information</u>) contribute to the diet in different proportions according to species. Littoral invertebrates (n = 21) constitute approximately 23–61% of perch diet (Figure 3a), 0–23% of pollan diet (Figure 3b), and 62–79% of roach diet (Figure 3c). The pollan diet as indicated by the mixing model is strongly influenced by calanoid zooplankton (19–62%) and to a lesser extent *Daphnia* spp. (1.4–33%). Percentage contributions of sources to diet listed in the text are measured at the 95% confidence level (data can be found in the <u>supplemental information</u>).

In short, as indicated by the stable isotope data (Figure 2), the mixing model suggests that pollan have less influence from littoral sources than perch, which have a more evenly mixed diet. Roach have the highest littoral contribution to their diet. However, the variation evident in the mixing model results, although due to source sample size issues, is evidence of the issues of stable isotope analysis and the clarity of conclusions that can be drawn from such data.

The sample size of stable isotope measurements of some sources was small, but the sample size of ¹⁴C measurements for these sources was smaller still (see <u>supplemental information</u>). Large ranges of Δ^{14} C values within a consumer/prey cohort lead to high standard deviations, again affecting the reliability of the model; consequently, the model results are not presented here. However, Figure 4 gives an indication of the differences in the dietary information provided by the utilization of Δ^{14} C measurements not shown by stable isotope analysis. Error bars represent the statistical uncertainty of Δ^{14} C measurements. There is no correlation between δ^{13} C and Δ^{14} C ($R^2 = 0.05$, p = 0.0885). The variation in species suggests there is no relationship between collection location and individual species Δ^{14} C value (ANOVA, F = 0.765, p = 0.3856).



Figure 3 (a) Proportions of dietary sources contributing to the diets of perch (*Perca fluvi-atilis*), (b) pollan (*Coreganus autumnalis*), and (c) roach (*Rutilis rutilis*). Mixing models were generated using SIAR v 4.2, R v 2.12.1 and are described in more detail in the text. Shaded gray levels in plot refer to 5, 25, 75, and 95% confidence intervals.

Calanoid Zooplankton

Daphnia

Chironomidae

Littoral Invertebrates

It is clear from ¹⁴C measurements that individual fish from all species utilize divergent carbon sources. Δ^{14} C of the majority of fish are enriched relative to summer calanoid Δ^{14} C. Three small perch (9, 9.3, 9.4 cm in length) have the most depleted Δ^{14} C (-68.2‰, -52.7‰, and -60.8‰, respectively).



Figure 4 $\Delta^{I4}C$ and stable carbon isotope measurements from Lower Lough Erne fish. Error bars represent the uncertainty on measurements. Horizontal lines denote the mean $\Delta^{I4}C$ as measured from calanoid zooplankton in winter and summer. Winter calanoid $\Delta^{I4}C$ results are analogous to those of modern terrestrial carbon, summer to autochthonous algal $\Delta^{I4}C$.



Figure 5 Stable isotope values of Lower Lough Erne pelagic invertebrate species. Points with error bars denote mean algal and POM isotope values with their associated standard deviations.

Invertebrate Diet

Stable isotope analysis of Lower Lough Erne invertebrates demonstrates a distinction between groups of invertebrates (Figure 5). ¹⁵N values (winter = $23.4 \pm 1.2\%$, summer = $16.6 \pm 3.7\%$) of the calanoid zooplankton *Eudiaptomus gracilis* samples in winter are highly enriched relative to phytoplankton, POM, and the other components of the food web including predatory fish. While they drop in summer, they are still higher than expected from an algal diet. *E. gracilis* is an omnivorous species that preferentially feeds on phytoplankton (Maguire and Grey 2006; Ger et al. 2011). Their δ^{13} C results (winter = $-32.4 \pm 2.5\%$, summer = $-33.4 \pm 2.4\%$) are inconclusive; the results could describe an algal (δ^{13} C = $-33.4 \pm 3.8\%$, δ^{15} N = $+9.3 \pm 4.7\%$) or particulate organic matter (δ^{13} C = $-31.4 \pm 1.3\%$, δ^{15} N = $+13.4 \pm 7.3\%$) carbon source. Their Δ^{14} C values are more clear: winter values are enriched and within the range of lake DOC values (mean winter calanoid Δ^{14} C = $-4.0 \pm 6.1\%$). However, the summer results are within range of algal sample Δ^{14} C values (Δ^{14} C = $-50 \pm 4.1\%$).



Figure 6 Invertebrate Δ^{14} C and δ^{13} C measured from selected Lower Lough Erne invertebrate species. Winter samples were collected in February and March, and summer samples in May and August 2011. Points with error bars denote mean algal and POM isotope values with their associated standard deviations. Mean DOC Δ^{14} C (relatively consistent over the duration of the study) is shown as a horizontal line; δ^{13} C values were unreliable for DOC samples. Dashed lines denote the error associated with the mean.

Daphnia spp. ¹⁵N is also enriched but more so in the summer (δ^{15} N in winter = +12.8 ± 1.7‰, summer = +17.5‰) despite increased algal abundance. *Daphnia* Δ^{14} C values (Figure 6) are depleted (Δ^{14} C = -108.1 ± 5.2 to -91.9 ± 3.6‰) and are within the range of subfossil POM values (Δ^{14} C = -83.1 ± 4.0 to -191.9 ± 12.4‰).

DISCUSSION

The results from Lower Lough Erne samples show evidence that three dietary carbon sources exist in the lake.

Autochthonous Fraction

Carbon partially derived from bedrock weathering is utilized in primary production, leading to a mean algal value of $\Delta^{14}C = -58.8 \pm 9\%$ (excluding the May cyanobacteria sample). This ¹⁴C depletion results from the freshwater reservoir offset and is linked to water alkalinity (Keaveney and Reimer 2012).

Modern/Recent Radiocarbon

Riverine DOC sampled in October 2011 had a positive bomb ¹⁴C signal ($\Delta^{14}C = +51.9 \pm 2.9\%$), indicating a modern or recent photosynthetic carbon source (Levin et al. 2013).

Labile Allochthonous Carbon

Lake DOC is ¹⁴C-enriched ($\Delta^{14}C = -17.4 \pm 19.4\%$) relative to DIC and POC in the lake, yet depleted relative to modern values. It is consistent with that from incoming River Bannagh in winter, indicating a terrestrial source.

Subfossil Allochthonous Carbon

POM Δ^{14} C values are depleted relative to anything else in the lake, indicating the presence of a subfossil terrestrial carbon source in the lake (Δ^{14} C = -122.3 ± 61.3‰).

Invertebrate Food Web

While the sample size of invertebrate prey was small, there is clear evidence that terrestrial support of the food web is extensive. Invertebrates utilize both energetically-rich autochthonous and terrestrial carbon of lower nutritional value. Within the terrestrial carbon pool, both subfossil, detrital carbon with low nutrient value and chemically labile carbon are consumed by different invertebrates.

The highly positive δ^{15} N values of calanoid zooplankton indicate the presence of a microbial loop with bacteria metabolizing terrestrial carbon, which is then consumed by other species (Meyer 1994). The existence of a microbial loop is an indicator of a terrestrial diet, but stable isotope analysis does not indicate whether the terrestrial source utilized by Lower Lough Erne calanoid zooplankton is detrital, or whether it is a labile terrestrial carbon source.

 Δ^{14} C values of calanoid zooplankton measured in winter and summer demonstrate the significant differences in carbon sources utilized (Figure 6). Δ^{14} C measurements from winter calanoid zooplankton (winter = -7.5 ± 14.7‰) are within the range of lake DOC Δ^{14} C (-4 to -40‰, Figure 6). In the summer, calanoid Δ^{14} C values are comparable to algal Δ^{14} C, showing a clear dietary switch with algal blooms. The calanoid invertebrates are clearly selecting bio-energetically richer carbon, labile terrestrial carbon in the winter with low algal abundance and a switch to an algal diet in the summer. They are not utilizing subfossil carbon sources, showing a clear selection against detrital carbon.

Daphnia Δ^{14} C values are more depleted than algal Δ^{14} C. Daphnia have been shown to feed on the algal portion of the POM (Rautio et al. 2011). However, Daphnia Δ^{14} C values indicate that they are consuming a subfossil carbon source that is older than the algal carbon pool, suggesting that they are also feeding on a detrital terrestrial portion of POM. The availability and biochemical quality of carbon seems to drive the dietary preferences of calanoid zooplankton, while Daphnia spp. derive carbon from particulate sources of detrital subfossil carbon.

Fish

The stable isotope values of fish are as expected; benthic/planktonic pollan have a strongly predatory signal and their isotope values reflect their specialized diets. Roach and perch, again as expected, have a more mixed littoral and pelagic diet, showing their less selective diet. The evidence for terrestrial subsidy of fish diet is not as clear when using stable isotope analysis alone as δ^{13} C overlaps with terrestrial values. The evidence for terrestrial support of all fish species is only found in the results from ¹⁴C analysis.

A range of Δ^{14} C values are evident in individual Lower Lough Erne fish. While stable isotopes can indicate that terrestrial carbon may be a component of diet of Lower Lough Erne invertebrates, Δ^{14} C clearly shows that individual fish within and between species can obtain carbon through different food web pathways including those supported by terrestrial carbon. Even Δ^{14} C values of open/deep water fish (e.g. pollan) are enriched, indicating that terrestrial carbon has been transferred to the highest trophic levels.

An individual pollan and roach each had positive Δ^{14} C values indicating the presence of bomb 14 C in their diet. No invertebrate species measured in this study had an equivalent bomb 14 C signal. However, DOC from a tributary river to Lower Lough Erne, the River Bannagh (see Figure 1), collected in October 2011, had a highly positive Δ^{14} C (+51 ± 2.9‰), which corresponds to atmospheric 14 C from 2005–2009 (Hua and Barbetti 2004; Levin and Kromer 2004; Levin et al. 2008, 2013), thus representing modern/recent photosynthetic carbon from leaf fall, grass, or riparian plant material.

A modern terrestrial carbon source is clearly a part of the food web of fish with positive Δ^{14} C values, although the mechanism by which this carbon enters the food web is unclear. Enriched bomb Δ^{14} C as seen in Lower Lough Erne can only be obtained from the products of a terrestrial plant or its consumer. Much of the shoreline of the lake consists of shrubs or trees, while the more sheltered bays have extensive beds of emergent macrophytes, which could be contributing a source of modern carbon via leaf fall and plant senescence. The enriched Δ^{14} C values of the individual pollan and roach show they must be utilizing prey consuming this modern source of terrestrial material in addition to autochthonous and terrestrial food sources. This prey source is missing from our analysis; we did not measure any source contributing to diet other than riverine DOC that had sufficient enrichment to cause these Δ^{14} C values.

The use of methane can result in ¹⁴C-enriched/depleted results depending on the source of methane (Billett et al. 2012; Garnett et al. 2013). However, samples in Lower Lough Erne are not sufficiently ¹³C-depleted to provide evidence for methanogenesis in the system (Pasche et al. 2011). The absence of an enriched Δ^{14} C bomb signal in the pelagic invertebrate species measured in this study suggests that there may be a benthic source of modern carbon that was not sampled in the current study.

Perch have a mixed pelagic and littoral diet according to the relatively large range of δ^{13} C values. The majority of perch Δ^{14} C measurements fall within the range of those of zooplankton with an algal diet such as summer calanoid samples. They do not appear to have a subfossil carbon source; all except three are enriched relative to mean calanoid zooplankton Δ^{14} C. The exceptions are the three oldest perch, which clearly have a prey source consuming a slightly older source of carbon, although Δ^{14} C indicates they are still within the range of autochthonous carbon sources. They were collected from K30 (Table 1), at a depth of 1–4 m. These perch are among the smallest and are likely to feed in the littoral zone, as indicated by the stable isotopes; Δ^{14} C may also indicate a littoral carbon source [values are within 2σ error of littoral zone food sources such as *Asselus aquaticus* (Δ^{14} C = $-57.3 \pm 3\%$) and biofilm (Δ^{14} C = $-56.6 \pm 12.2\%$)].

Roach are omnivorous and this is clearly shown in the stable isotope and ¹⁴C measurements. The range of Δ^{14} C values indicates they are consuming invertebrates feeding in terrestrial and/or autochthonous carbon pools. None are more depleted than algal Δ^{14} C, indicating that terrestrial carbon is more important in the food web of roach. One roach, like the pollan sample, is enriched with bomb ¹⁴C, indicating they are consuming prey utilizing a modern terrestrial carbon pool.

CONCLUSION

The use of the ¹⁴C method in analyzing the Lower Lough Erne food web has elucidated the sources of carbon utilized in the lake. Terrestrial support of the food web was expected given previous research and water chemistry (e.g. DOC loading, chlorophyll *a* values). The method allows for clear separation of modern/recent, labile, and subfossil terrestrial carbon, in addition to the partitioning of terrestrial and autochthonous carbon sources. The method also shows the extent to which the pelagic food web is driven by terrestrial carbon sources, to the highest trophic levels; terrestrial support is underestimated or overlooked with stable isotope analysis alone.

The distinction between modern, labile, and detrital subfossil terrestrial carbon in the food web made using ¹⁴C analysis is essential when analyzing carbon dynamics in freshwater lakes. The data have implications for biodiversity management as the proliferation of invasive species takes place. Terrestrial subsidies of freshwater lakes are likely to increase with climate and land-use changes; this information is of paramount importance in the overall management of freshwater areas and their catchments in the near future.

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