

MICADAS: ROUTINE AND HIGH-PRECISION RADIOCARBON DATING

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ABSTRACT. The prototype mini carbon dating system (MICADAS) at ETH Zurich has been in routine operation for almost 2 yr. Because of its simple and compact layout, setting up a radiocarbon measurement is fast and the system runs very reliably over days or even weeks without retuning. The stability of the instrument is responsible for the good performance in highest-precision measurements where results of single samples can be reproduced within less than 2‰. The measurements are described and the performance of MICADAS is demonstrated on measured data.

INTRODUCTION

The mini carbon dating system (MICADAS) at ETH Zurich was built 4 yr ago (Synal et al. 2007), based on principles already used at the 0.5MV system called TANDY (Synal et al. 2000). In the first 2 yr, MICADAS was mainly used for experimental development, while routine radiocarbon measurements were still performed on the large 6MV EN-tandem, as they had been successfully done for almost 30 yr (Suter et al. 1984a; Bonani et al. 1987). During the last 2 yr, we made a smooth transition of routine ¹⁴C measurements from the large accelerator to the MICADAS, mainly because the latter requires little maintenance and runs unattended. Here, we present our experience with the MICADAS in routine operation from the past 2 yr. We were also encouraged to test the MICADAS for highest-precision measurements, because already the first measurements had shown that it is possible to measure samples very reproducibly (Synal et al. 2007).

METHOD

Sample Preparation

Most samples were graphitized on a semi-automated system after sample cleaning and combustion in closed quartz tubes (Hajdas et al. 2004). The system was built 8 yr ago and produces about 2 mg of graphite deposited on 7 mg of iron. This is a rather large amount optimized for the measurement on the large 6MV tandem accelerator, which is equipped with a negative ion source where the samples are moved during sputtering to avoid crater formation (Bonani et al. 1987). On the other hand, only 1–1.5 mg of graphite are typically used for a measurement using MICADAS.

The wood samples for high-precision measurements as shown below were cleaned with a base-acid-base-acid bleaching procedure (Němec et al. 2010) and then graphitized on a recently developed system that is fully automated (Wacker et al. 2010b). The graphitization system is directly coupled to an elemental analyzer (EA) and runs without any user interaction after the samples are loaded into the EA and the iron catalyst is inserted into individual reactors. The amount of graphite has been adapted for the measurement on MICADAS (1 mg of graphite on 3.5 mg of Fe).

Some additional samples were also prepared at the Academy of Sciences in Heidelberg. These samples were tube combusted and graphitized on a semi-automated system, similar to the one at ETH

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(Unkel 2006). All samples were pressed pneumatically into a 1-mm hole of the sample holders with 250 MPa (equivalent to 200 N on a pin with 1 mm diameter).

Measurements

We have now measured more than 4000 solid samples and 1000 gaseous samples in the first 18 months of routine operation on the MICADAS. The machine runs very reliably for days without having any problems. MICADAS is designed to run unattended and features an automatic monitoring system. The hardware and the measured results are continuously checked, and in case of any anomaly, the automated measurement is stopped and/or the operator is informed by email or a short message on the mobile phone.

The main system parameters are given in Table 1. The source is equipped with a prototype sample changer that allows changing cassettes with 20 positions without interruption of the measurement. It is also possible to measure small CO₂ samples (<100 µg carbon) directly, without having to graphitize the samples (Ruff et al. 2007, 2010).

Table 1 Main parameters for MICADAS measurements.

Parameter	Value
Graphite	0.9 mg/2 mg ^a
Iron catalyst	3.5 mg/7 mg ^a
C ⁻ current (LE)	20–40 µA
Source potential	38 kV
Terminal voltage	195 kV
Stripper gas	N ₂
Transmission (¹² C)	43%
Measured ratio (1 F ¹⁴ C)	1.066 × 10 ⁻¹²
Precision (modern)	
routine	3–4‰
high-precision	1.5–2‰

^aDepending on the graphitization system, see text.

After turning on the ion source and the magnets, the system is left to stabilize for 2–3 hr. The tuning of the system takes about 30 min or less, even for highest-precision measurements. This is mainly because the only adjustable steering and focusing options for the ion beam are inside the ion source. The system is normally set up on Monday and is then left running without or only minor retuning over the rest of the week or even longer. Measurements were carried out with ¹²C currents between 20 and 40 µA on the low-energy side. The transmission (¹²C⁻/¹²C⁺) was always very stable at 43.1 ± 0.4% (based on all OXII standards measured during the first half of 2009; see Figure 1).

A typical ¹⁴C measurement on the MICADAS at ETH Zurich is done as follows: a cassette has 20 samples, including 3 standards and 2 blanks; all samples of a cassette are measured at least 5–6 times. A single measurement is subdivided into 10 cycles of 45 s each. For high-precision measurements, we used 4–6 standards and at least 3 blanks and the measurement time of 30–60 min per sample was increased to 2–3 hr by extending the cycle time to 60 s and by measuring each sample 10–15 times.

The recorded data set for each measurement consists of the number of ¹⁴C counts, the measurement time, high-energy currents of ¹²C⁺, ¹³C⁺, and also the ¹³C⁺ current of broken-up ¹³CH⁻ molecules (¹³C_H), injected together with ¹⁴C into the stripper (Suter et al. 1984b; Synal et al. 2007). The ¹⁴C

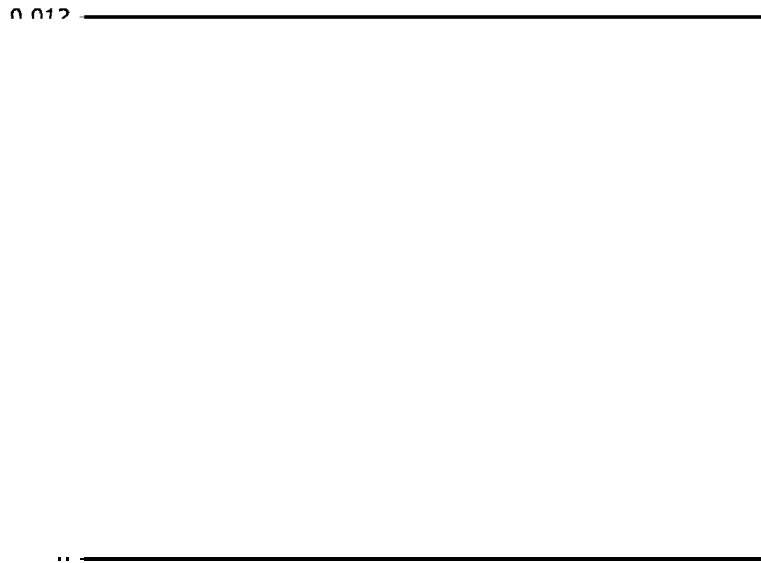


Figure 1 Correlation between the measured $^{14}\text{C}/^{12}\text{C}$ ratio and the ^{13}C current from the molecule breakup divided by the ^{12}C current for all blank samples measured in 2009 is shown. The slope of the plotted lines corresponds to our applied molecular correction ($450 \mu\text{A}^{-1}\text{s}^{-1}$, see text). An average blank sample lies on the solid line, and corresponds to a $^{14}\text{C}/^{12}\text{C}$ of 2.5×10^{-15} or 48,000 yr after molecular background correction. The scatter in the data comes mainly from the fact that the samples were prepared on 3 different graphitization systems over a relatively long time. For 1 set of samples prepared and measured together, the variation is significantly less.

counts were initially collected in a two-parameter ΔE -E spectrum gained using a gas ionization detector (Suter et al. 2007). Only the counts within a polygonal gate in the two-dimensional ΔE -E plot were accepted. Recently, we switched to a one-parameter (ΔE) measurement with a single channel analyzer (Schulze-König et al. 2010).

Data Analyses

Data analyses are performed with a computer program called BATS, described in detail by Wacker et al. (2010a). In short, the following steps were applied for the data reduction:

1. Molecular background subtraction, based on the measured $^{13}\text{C}_\text{H}$;
2. Blank subtraction;
3. Fractionation correction using the measured $^{13}\text{C}/^{12}\text{C}$ ($\delta^{13}\text{C}$) of the corresponding $^{14}\text{C}/^{12}\text{C}$ ratio;
4. Standard normalization using all standards.

The program also visualizes the data and verifies the quality of the acquired data with statistical tests. It also takes into account an additional uncertainty, reflecting the reproducibility of the results of the individual samples, and combines it with the uncertainty due to counting statistics, standard normalization, and blank correction. BATS is used to evaluate the data even while measurements are under way. The program evaluates the data without any user input within a few seconds and shows immediately if the desired precision has been reached and the measurement can be stopped.

DISCUSSION

Background and Blank

In the gas ionization detector, we observe a molecular background of $^{12}\text{C}^+$ and $^{13}\text{C}^+$ corresponding to typically to 0.004 to 0.006 fraction modern carbon ($F^{14}\text{C}$), which is about twice the real $^{14}\text{C}^+$ of a typical processing blank. Although we have installed a 2-parameter gas ionization detector system (Suter et al. 2007) on MICADAS, we cannot separate the background from real ^{14}C counts. Only a slight shift of the fitted peak center can be observed, which can only be used for diagnostics and not for data evaluation of real measurements.

The background derives from molecules with mass 14 ($^{13}\text{CH}^-$ or $^{12}\text{CH}_2^-$) that are broken up in the stripper. A charge exchange from 2+ to 1+ in the second accelerator part is then likely responsible for the additional gained energy of the ^{13}C or ^{12}C ions, which is required to pass the electrostatic analyzer. The energy measured in the gas detector confirms this assumption. An additional angular scatter is possibly responsible that the ions first also pass the magnet, which in principle separates ions of different momenta.

The molecular background in the detector is dominated by ^{13}C . However, ^{12}C can also get through the high-energy filter, which was confirmed by a measurement of graphite reduced with deuterium instead of normal hydrogen. In this case, ^{12}C is injected at the low-energy side as $^{12}\text{CD}^-$, which can be estimated to be about 100 times more abundant than $^{12}\text{CH}_2^-$ (based on observed $^{12}\text{CH}^-/^{12}\text{CH}_2^-$ ratios at the low-energy side). The result was a 50× increased molecular background in the detector (0.3 $F^{14}\text{C}$). However, the $^{12}\text{C}^+$ current measured at the high-energy side during mass-14 injection was about 10 times lower than the $^{13}\text{C}^+$ current. The molecular background therefore scales with the measured $^{13}\text{C}^+$ current from broken-up $^{13}\text{CH}^-$ molecules (Synal et al. 2007). This $^{13}\text{C}^+$ current is measured in a separate Faraday cup between the ^{12}C and the ^{13}C cup on the high-energy side (Synal et al. 2000). The correlation between this normalized molecular current ($^{13}\text{C}_\text{H}/^{12}\text{C}$) and the $^{14}\text{C}/^{12}\text{C}$ is shown in Figure 2 to be stable over time.

Typically, we see that the molecular background-corrected blanks of 0.002–0.003 $F^{14}\text{C}$ show a standard deviation of 0.0003 $F^{14}\text{C}$ for high-precision measurements (with long data acquisition time) and 0.0005 $F^{14}\text{C}$ for normal samples. The reason for this difference is that high-precision samples are always measured directly after preparation, whereas normal samples may be stored for up to 2 weeks, which seems to influence the blank variability. Additionally, any surface contamination from sample pressing has less influence on a long measurement compared to a short one. The measurement of 50,000-yr-old samples is possible with the background correction, whereas without this correction, the limit is 45,000 yr (calculated as twice the blank variability).

Stability and Reproducibility

The tuning of MICADAS for ^{14}C measurements is straightforward and once the system is tuned, it is very stable over time and rarely needs any retuning. The high stability is demonstrated in Figure 2, where the measurements of all OXII standards over a time period of 19 days are plotted (typically about 4 fresh OXII standards were mounted in cassette). During this time, the MICADAS system was not retuned. On the 16th and the 17th day, a gas measurement was performed and afterwards the same settings we had before for solid samples were reloaded (no tuning!).

The data points in the top pane (1) of Figure 2 show the measured raw ratio of $^{14}\text{C}/^{12}\text{C}$ with a scatter of 7.4‰, which is significantly higher than what can be expected from the counting statistics (4.3‰). The reason for this discrepancy lies mainly in the isotopic fractionation of the samples. The graphite produced in the individual graphitization system can change up to 10‰ in the $^{13}\text{C}/^{12}\text{C}$ ratio

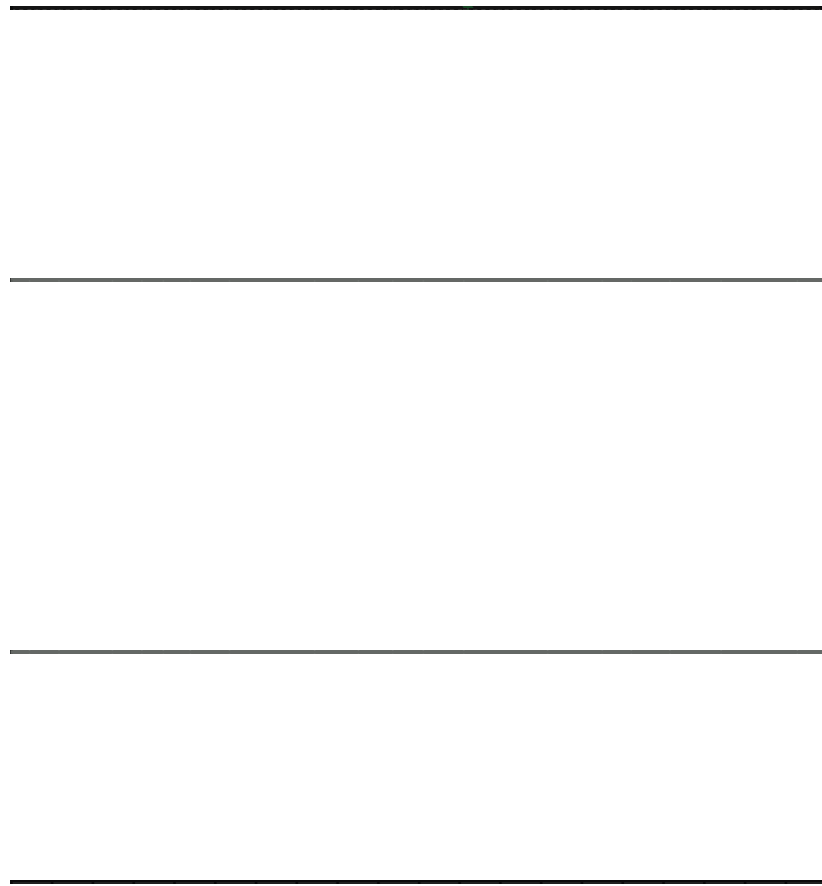


Figure 2 Raw $^{14}\text{C}/^{12}\text{C}$ ratios over time of single measurements of the OXII standard are shown in the top pane (1). In the second pane, the $^{14}\text{C}/^{12}\text{C}$ ratio corrected for the $^{13}\text{C}/^{12}\text{C}$ ratio is shown, once for the ^{14}C counts accepted by the single channel analyzer (2) and once the ^{14}C counts in the gate of the 2D-spectra from the gas ionization detector (3). The measured $^{13}\text{C}/^{12}\text{C}$ ratio is shown in the lower pane (4). The vertical lines separate the measurements of a set of samples in 1 cassette.

between samples and thus about twice as much for the $^{14}\text{C}/^{12}\text{C}$ ratio. If the $^{14}\text{C}/^{12}\text{C}$ ratios are then corrected for fractionation with the measured $^{13}\text{C}/^{12}\text{C}$ (shown in the lower panel (4) of Figure 1) as shown in the middle panel (3) of Figure 1, the scatter is 4.7‰, only slightly higher than expected from counting statistics. This demonstrates that MICADAS is extremely stable over days without any drifts in the ratios, and that it is possible to reliably correct for any fractionation from the sample preparation.

The bottom data points in the middle panel (3) of Figure 2 show also the $^{13}\text{C}/^{12}\text{C}$ corrected $^{14}\text{C}/^{12}\text{C}$ ratio. However, here the two-dimensional gating for the ^{14}C particle identification was used instead of gating with the single channel analyzer. It is obvious that a small trend over days is now visible and ~3% ^{14}C counts are lost, because the residual energy signal may drop below the low level discriminator that cuts off the noise. We ceased using the two-dimensional gating because the positive influence on the background level was only minor (−5%) while the simple gating with the single channel analyzer resulted in an increased yield (+3%) and better long-term stability (see Figure 2).

The cassettes 7, 12, and 16 shown in Figure 2 were high-precision measurements where 20 samples were measured over 2 days to about 700,000 ^{14}C counts. These samples were graphitized on the latest fully automated graphitization system (Wacker et al. 2010b) and tend to show much lower variation in fractionation than other samples prepared on a semi-automated system (all other cassettes).

The performance in routine analyses is given in Figure 3 on measurements of IAEA-C5 and IAEA-C3 samples. The mean measured value for the IAEA-C3 of (0.2301 ± 0.00034) F^{14}C is in good

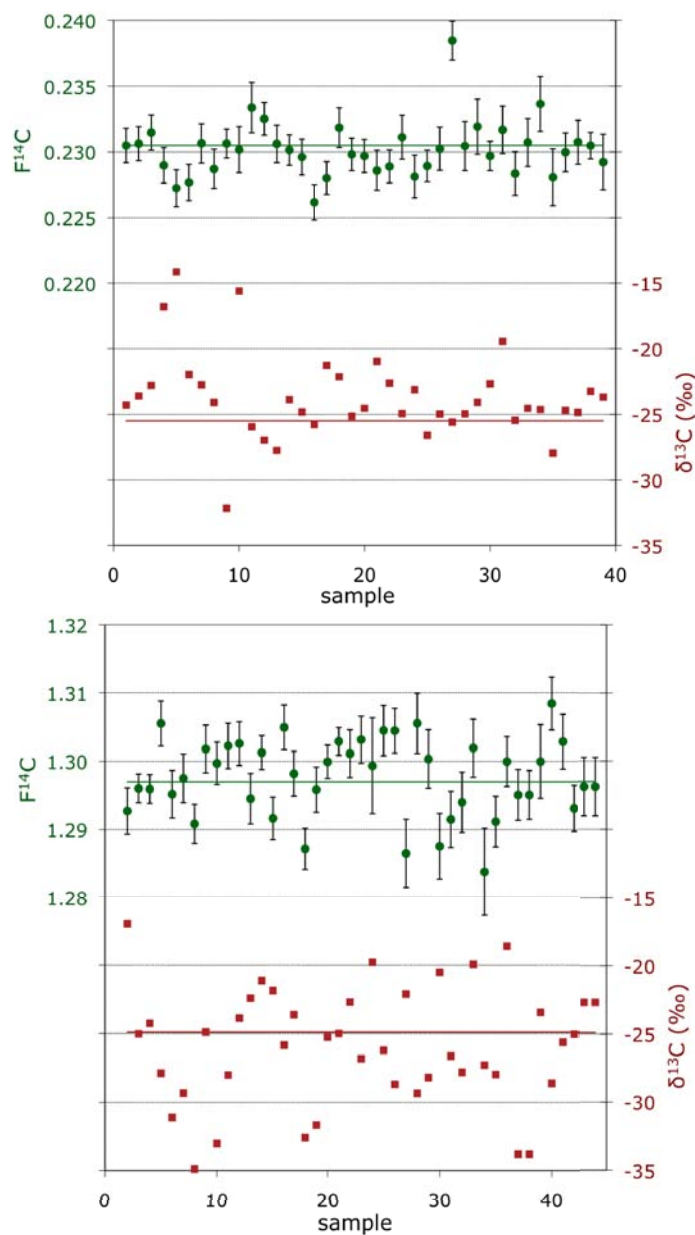


Figure 3 Measurements of the reference materials IAEA-C5 and IAEA-C3 during routine operation are presented. The solid line represents the reference value. The full circles on top are the F^{14}C values and the squares below are the corresponding $\delta^{13}\text{C}$ values.

agreement with the reference value of $(0.2305 \pm 0.0002) F^{14}C$. The standard deviation of these measurements is with $0.0021 F^{14}C$ only slightly higher than the mean uncertainty from counting statistics and the blank correction for a single measurement ($0.0015 F^{14}C$). Also, the IAEA-C3 value of $(1.298 \pm 0.001) F^{14}C$ is in perfect agreement with the reference value of $(1.297 \pm 0.001) F^{14}C$. Here, the standard deviation of the measurements ($0.006 F^{14}C$) is significantly higher than the uncertainty from the counting statistics and the blank correction of $0.004 F^{14}C$. We therefore conclude that for our routine measurements, we have an additional unknown uncertainty for the sample reproducibility (external uncertainty) of 3%. This uncertainty is added to all our routine measurements (Wacker et al. 2010a).

Precision

The MICADAS performs very well in high-precision measurements. Figure 4 shows the measured ratios of the 4 OXII standards we use for normalization of a set of samples. These standards were all measured to about 1‰ counting statistics. The presented fractionation-corrected $^{14}C/^{12}C$ ratios show a distribution that is basically defined by counting statistics. All high-precision measurements indicate that the theoretical limit for the reproducibility of the standards is well below 1‰. At the beginning of a measurement, a slight surface contamination (increased blank) sometimes causes a small additional scatter in the measurements of a single sample, but high-precision measurements over long time periods seem to have less scatter in addition to the counting statistics of ^{14}C . This means an increased measurement time not only shows a higher counting statistics, but also seems to be more reliable.

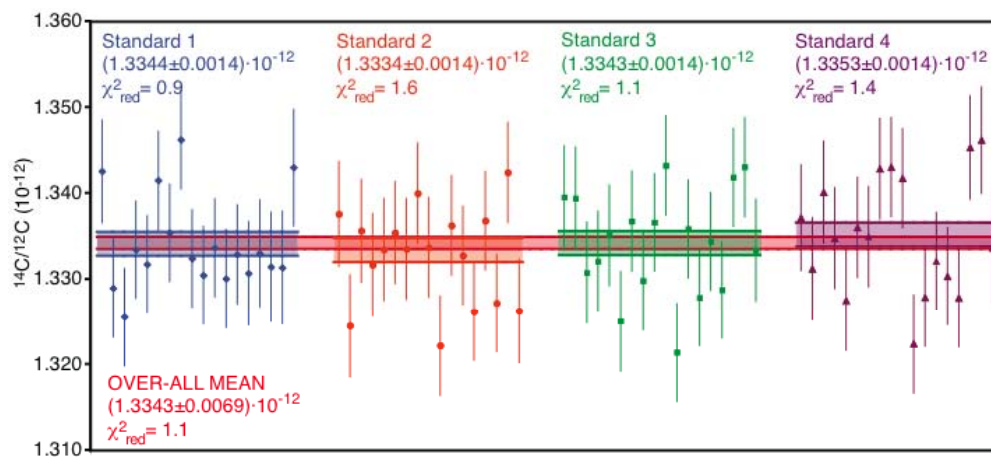


Figure 4 $^{14}C/^{12}C$ ratios of 4 OXII standards (17 passes) in a high precision measurement are shown. Each standard is measured to $\sim 1\text{‰}$ counting statistics. The bars indicate the $1\text{-}\sigma$ range of the corresponding mean values.

The reproducibility of high-precision measurements on 2 real samples is given in Table 2. Both samples, a parchment and a seal cord from the Middle Ages, were each cleaned twice and graphitized 4 or 5 times, respectively, on the 2 different graphitization systems at ETH (Hajdas et al. 2004; Wacker et al. 2010b). Subsamples of the 2 samples show 2 distinct ages with a perfect repeatability within their final uncertainties of ~ 20 yr (2.5‰). This suggests that the subsamples of each sample can be averaged with an uncertainty of about 10 or 8 yr ($\sim 1\text{‰}$), respectively. The uncertainties of these mean values are lower than that of the calibration curve IntCal04 (Reimer et al. 2004), which has an uncertainty of 12–13 yr in ^{14}C ages in this time range (with a decadal resolution in the samples entering IntCal04).

Table 2 List of the single measurements of the parchment and the seal cord with their mean values, given in ^{14}C yr BP with $\pm 1\text{-}\sigma$ uncertainties. The $\delta^{13}\text{C}$ is the measured isotopic ratio of the prepared graphite and not necessarily of the original sample.

Lab number	Sample type	^{14}C age (yr BP)	$\delta^{13}\text{C}$ (‰)
ETH-36716.1	parchment	888 ± 20	-22.6 ± 1.1
ETH-36716.2	parchment	878 ± 19	-20.7 ± 1.1
ETH-36716.3	parchment	882 ± 19	-23.9 ± 1.1
ETH-36716.4	parchment	875 ± 19	-22.3 ± 1.1
ETH-36716	parchment	881 ± 10	-22.4 ± 0.6
ETH-36717.1	seal cord	800 ± 20	-24.0 ± 1.1
ETH-36717.2	seal cord	808 ± 19	-29.1 ± 1.1
ETH-36717.3	seal cord	833 ± 18	-25.5 ± 1.1
ETH-36717.4	seal cord	808 ± 18	-27.1 ± 1.1
ETH-36717.5	seal cord	800 ± 17	-27.7 ± 1.1
ETH-36717	seal cord	809 ± 8	-26.7 ± 0.9

Therefore, it was decided to remeasure dendrochronologically dated tree-ring samples of the southern German oak chronology (Friedrich et al. 2004) in the time range of 1000 to 1200 calendar yr BP; the results are given in Table 3. The samples, covering 5 yr each, were measured to a precision of about 15 yr (2‰), including an uncertainty for the sample preparation of 0.8‰. The sample-to-sample ^{14}C age variability (26 yr) is very low and compares very well with the decadal IntCal04 raw data in the same time range from the different labs (Pearson and Stuiver 1986; Stuiver et al. 1998; Hogg et al. 2002), which show a variability between 30 and 35 yr from sample to sample in 10-yr resolution. In principle, with our 5-yr resolution we are sampling some of the production signal caused by the 11-yr solar cycle (Stuiver and Braziunas 1993), hence the lower sample-to-sample variability in our data is unexpected.

Table 3 Results of measurements on 5-yr tree-ring samples of 2 oaks from Eichstaett, southern Germany (Eichstaett 18, AD 1111–1269; Eichstaett 9, AD 1271–1305). Uncertainties include both counting statistics and sample preparation.

Lab number	Age range yr AD	^{14}C age (yr BP)
ETH-37133	1111–1115	968 ± 15
ETH-37134	1116–1120	959 ± 15
ETH-37135	1121–1125	955 ± 15
ETH-37136	1126–1130	939 ± 15
ETH-37137	1131–1135	954 ± 15
ETH-37138	1136–1140	966 ± 14
ETH-37139	1141–1145	959 ± 14
ETH-37140	1146–1150	960 ± 14
ETH-37141	1151–1155	926 ± 15
ETH-37142	1156–1160	927 ± 15
ETH-37143	1161–1165	895 ± 15
ETH-37144	1166–1170	899 ± 14
ETH-37145	1171–1175	909 ± 15
ETH-37146.1	1176–1180	945 ± 15
ETH-37146.2	1176–1180	920 ± 14
ETH-37147.1	1181–1185	909 ± 15
ETH-37147.2	1181–1185	868 ± 14
ETH-37148	1186–1190	882 ± 14
ETH-37149	1191–1195	899 ± 14

Table 3 Results of measurements on 5-yr tree-ring samples of 2 oaks from Eichstaett, southern Germany (Eichstaett 18, AD 1111–1269; Eichstaett 9, AD 1271–1305). Uncertainties include both counting statistics and sample preparation. (*Continued*)

Lab number	Age range yr AD	¹⁴ C age (yr BP)
ETH-37150	1196–1200	909 ± 14
ETH-37151	1201–1205	909 ± 15
ETH-37152	1206–1210	885 ± 14
ETH-37153	1211–1215	849 ± 14
ETH-37154	1216–1220	899 ± 14
ETH-37155.1	1221–1225	867 ± 15
ETH-37155.2	1221–1225	869 ± 14
ETH-37156.1	1226–1230	875 ± 14
ETH-37156.2	1226–1230	875 ± 14
ETH-37157.1	1231–1235	807 ± 14
ETH-37157.2	1231–1235	829 ± 15
ETH-37158.1	1236–1240	863 ± 14
ETH-37158.2	1236–1240	843 ± 15
ETH-37159.1	1241–1245	819 ± 14
ETH-37159.2	1241–1245	805 ± 15
ETH-37160	1246–1250	837 ± 14
ETH-37161	1251–1255	821 ± 15
ETH-37162	1256–1260	813 ± 17
ETH-37163	1261–1265	790 ± 15
ETH-37164	1266–1269	775 ± 14
ETH-37165	1271–1275	748 ± 14
ETH-37166	1276–1280	757 ± 14
ETH-37167	1281–1285	723 ± 14
ETH-37168	1286–1290	734 ± 14
ETH-37169	1291–1295	704 ± 14
ETH-37170	1296–1300	705 ± 14
ETH-37171	1301–1305	702 ± 14

Though the data set looks very consistent, we see on average a significant offset of 26 yr (~3‰) to the calibration curve (see Figure 5). At present, we cannot explain this offset as all our measurements of secondary standards agree fully with the consensus values (see e.g. Figure 3). We do not know if the offset originates from systematic effects of our measurement procedures. Some of the wood samples presented were also treated with a simple acid-base-acid procedure for comparison and do not show the 26-yr offset (Němec et al. 2010). The applied rigorous pretreatment with base-acid-base-acid-bleaching could therefore be responsible for the observed offset.

CONCLUSIONS AND OUTLOOK

The operation of the mini carbon dating system MICADAS is extremely stable over time and no user interaction is required over several days. The stability of the measurements over time allows us to make even highest-precision measurements that until recently were a niche for gas counters (Kromer and Münnich 1992) and advanced liquid scintillation facilities. It goes beyond high-precision measurements on large accelerators which suggest that modern samples can be measured down to a precision of 2‰ (Graven et al. 2007; Meijer et al. 2006).

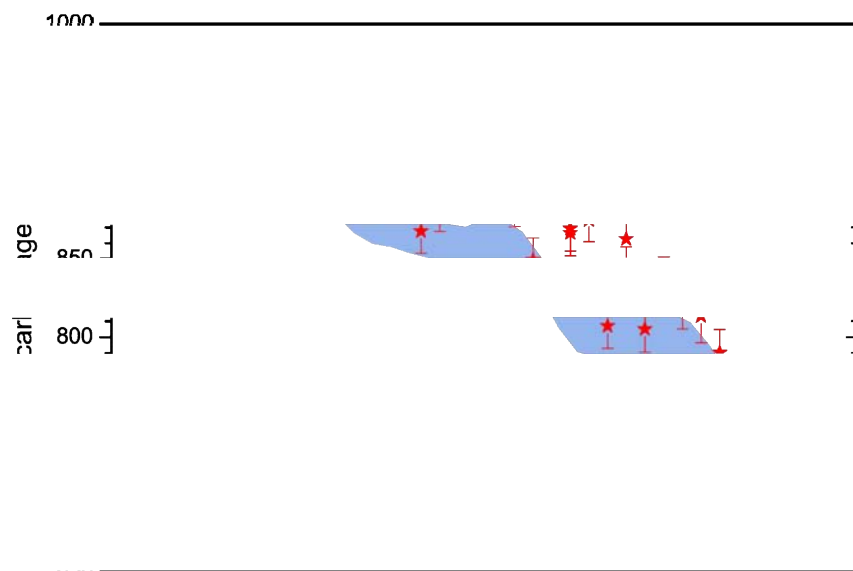


Figure 5 The 5-yr oak tree-ring samples from southern Germany samples measured at ETH Zurich are plotted with $1\text{-}\sigma$ uncertainties together with the IntCal04 (Reimer et al. 2004) calibration band (also $1\text{-}\sigma$). The sample-to-sample variation is only 3‰; however, an offset of about 3‰ can be observed.

As a word of caution, we have seen a significant offset in our data when remeasuring the calibration curve. This discrepancy will be investigated further to clarify whether this offset is a yet unknown systematic effect of our measurement procedures or whether it depends on the applied pretreatment.

Improvements are still possible and desirable. For example, high-precision measurements require high counting statistics and hence need a long measurement time. Higher ion currents from the ion source are thus desired. Therefore, we are now replacing the prototype ion source from the MICA-DAS with a modified source, successfully used on the BioMICADAS (Schulze-König et al. 2010). This should allow us to run the source with 2–3 times higher C^- currents.

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