Heat Effects on Nutrient Release from Soils Under Ponderosa Pine

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Highlight: Litter and mor of ponderosa pine-forest soils released more watersoluble K and P following heating to 200 C than at higher or lower temperatures. The upper A1 horizon released the most water-soluble P and K following, respectively, 200 C and 500 C heat treatments. Total nitrogen decreased in the samples heated in excess of 200 C. Prescribed burning to control noncommercial pine on rangeland apparently would have little detrimental effect on K and P availability and cause a slight reduction in total nitrogen if heating is kept below 200 or 300 C.

Wells (1971, p. 90) suggested a need to "...develop methods to predict the results of a burning treatment." Prescribed burning affects the soil by destroying, charring, or heating the surface litter, the underlying partially decomposed mor or O-horizon, and the upper mineral soil layer. The effect of heat on nutrient release from soils underlying ponderosa pine (*Pinus ponderosa* Laws.) was studied in the laboratory. Soils were collected from the forest-grassland ecotone in the foothills of the Black Hills in South Dakota.

Literature Review

Effects of prescribed burning on soil physical and chemical properties and on fertility were summarized by Ralston and Hatchell (1971) and by Wells (1971). In general, the combustible organic surface layer is affected most and may be partially or entirely destroyed. High temperatures may occur in mineral soil surfaces if the fuel supply is large and burned to "white ash." However, temperature decreases rapidly with depth in the soil and with time following the fire.

Philpot (1970) studied the pyrolysis of pine needles and reported the maximum volatization was between 175 to 350 C with endothermic and exothermic reactions occurring, respectively, at 210, 350, and 400 C and at 300 and 475 C. At 400 C, 35% of the organic compounds remained in the ash. Hosking (1938) reported soil organic matter lost some weight below 100 C and then underwent nondestructive and destructive distillation, respectively, from 100 to 200 C and from 200 to 300 C. In the 200 to 300 C range, loss of organic matter increased to about 85% of the total. Ignition of the carbonaceous residue occurred above 300 C. Thus, the maximum change in organic matter would probably occur below 300 C. Additional changes could occur from reactions of the inorganic elements released from the organic matter and from the soil, expecially at higher temperatures.

Many inorganic soil materials, particularly the partially weathered micaceous minerals and clays, dehydrate at relatively low temperatures. Dehydration, in addition to an inversion of low temperature crystalline forms to higher temperature forms, could increase or decrease the solubility of P and K from the mineral. At higher temperatures, chemical changes of the mineral surfaces could be caused by alkaline or alkaline-earth compounds, either from the heated minerals or from organic matter combustion. Such changes might increase or decrease the solubility of P or K depending upon the chemical compounds formed when the material cools. Rapid heating and cooling may break a mineral apart as it expands and contracts. The fresh unweathered surfaces could release P and K more rapidly than weathered surfaces.

Materials and Methods

Duplicate samples of litter and A1 horizon were collected at 6 locations at the east (locations 1 through 4) and north (locations 5 and 6) foothills of the Black Hills where the ponderosa pine encroaches onto the plains grassland. Mor samples also were collected, except at location 5 where none was present. Litter consists primarily of partially decomposed, but identifiable, pine needles. Mor consists of almost completely decomposed organic debris beneath the litter. The A1 horizon material consisted of the upper 4 cm of the mineral soil. Soils were derived from Dakota sandstone (locations 1, 2, and 3), from acid to weakly calcareous Graneros shale (location 4), from very calcareous Spearfish shale (location 5), and probably from Dakota sandstone (location 6).

Samples reported for the 25 C heat treatment were at laboratory temperatures; the 50, 100, and 200 C samples were heated in a laboratory drying oven; and the 300, 400, and 500 C samples were heated in a muffle furnace. These temperatures span the range found in a controlled burn of dense second-growth ponderosa pine conducted near one of the sample sites. Very thin, rectangular plates were striped with several crayons that melted at different temperatures. The plates were wrapped with aluminum foil and inserted into the litter, mor, and the upper 6 cm of the mineral soil.

For water-soluble P and K determinations, 3 to 5 gram samples were heated for 1 hour at the designated temperature, cooled, transferred to 125 ml Erlenmeyer flasks, extracted with 100 ml of H₂O for 10 minutes on an oscillating shaker, and the supernatant was filtered through Whatman No. 42 filter paper. K was determined with a flame photometer and P with the chloromoybdic acid-colorimetric method (Jackson, 1958). Total water-soluble P (includes organic-bound P) in the leachate and total P and K of the samples were determined following a perchloric acid digestion. Separate samples were used for the total N determination by the Kjeldahl procedure.

Results and Discussion

The nitrogen content of the litter, mor, and A1 horizon samples (Fig. 1) did not decrease until combustion was initiated at temperatures greater than 200 C. About one-half to two-thirds of the nitrogen was lost when samples were heated to 400 C, which corresponds to organic matter or nitrogen losses reported by Philpot (1970) and DeBell and Ralston (1970); and only a trace remained at 500 C (determined, but not reported in Fig. 1).

Nitrogen loss depended on the degree of combustion of the sample. At 200 C, the samples were darkened, but charring

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was slight. The quantity of white ash in the charred material increased from 300 to 500 C. At 300 C, some plant fragments burned more completely than others; therefore, grayish-ash fibers were mixed with easily identified charred plant fragments. Presumably, combustion was more complete where the grayish ash was formed and the temperature was locally greater than the furnace temperature. Because of the probable sample and furnace temperature differences, the 300 C heat treatment was not used on samples for water- soluble K and P extraction.

Heat effects on the release of total and

Table 1. Approximate total K and P contents (ppm) of unheated (25 C) and 500 C-heat-treatment samples and water-soluble K and P from 200 C-heat-treatment samples.

		Total K		Water-soluble K	Total P		Water-soluble P	
Location	Sample ¹	25 C	500 C	200 C	25 C	500 C	200 C	
1	Litter	1531	1322	325	548	551	111	
	Mor	1265	1265	216	187	178	27	
	A1 horizon	833	833	66	89	85	4	
4	Litter	1965	1898	458	822	823	148	
	Mor	4333	5400	333	794	833	70	
	A1 horizon	4766	5833	142	664	765	26	
5	Litter	2433	2733	1024	544	560	109	
-	A1 horizon	4666	5699	92	533	542	29	

¹Single sample analyzed from one of the duplicate samples at the locations. Samples were heated for 1 hr.



Fig. 1. Effects of heat treatments on the total N and water-soluble P and K contents of litter, mor, and upper A1 horizon of Ponderosa-pine-forest soils.

water-soluble K from the samples are more complex than those for N. Total K contents of some samples heated to 500 C were equal or greater than the unheated samples (Table 1); volatilization of K (Jackson, 1958, p. 334) may not occur. Lower total K content of the unheated samples (compared to heated samples) was probably caused by incomplete digestion of the unground unheated samples in the perchloric acid. This digestion is recommended for finely ground samples. Thermal fracturing of K-rich minerals could substitute for grinding so that extraction would be greater from heated than from unheated samples. Quartz in the sandstone-derived soils at location 1 contains little K, so thermal fracturing would not increase K release (Table 1). Shale-derived soils at locations 4 and 5 released more K following 500 C heating than the unheated samples. Presumably, the K was released by thermal fracturing of micaceous minerals (Raman and Jackson, 1965). A loss of K by volatilization may be compensated for by additional release of K from minerals following heating. In any event, relative amounts of the water-soluble K extracted from litter, mor, or A1 horizon after heating to different temperatures were similar for the different locations.

The water-soluble K released was largest from litter and mor heated at 200 C. K may have been released from the partial decomposition and rearrangement of chemical bonds of the organic matter. At temperatures greater than 200 C, K may have been volatilized or combined into insoluble compounds with silicate and other inorganic anions released from the organic matter. Regardless of the cause, heating from 200 to 500 C decreased the water-soluble K content of the litter and mor. Heating to temperatures greater than 200 C increased the water-soluble K extracted from the A1 horizon samples at locations 4, 5, and 6 and not from the other 3 locations. Thus,

Table 2. The water-extracted P (%) in the PO₄-P form at each heat treatment for the profile samples. Values are averages of all analyses from the seven locations.

	Temperature of heating ¹								
Sample	25 C	50 C	100 C	200 C	400 C	500 C			
Litter	71.85	78.30	74.85	37.23	91.31	53.91			
Mor	96.30	113.72	103.08	39.70	80.78	67.68			
A1 horizon	26.20	33.93	36.93	101.87	92.18	94.60			

¹All samples heated for 1 hr.

release of water-soluble K by thermal fracturing in A1 horizon samples from locations 4, 5, and 6 apparently is greater than the loss from volatilization or fixation in insoluble forms.

Phosphorus is less apt to volatilize than K at high temperatures. Differences in total P contents of 25 C samples and those heated to 500 C (Table 1) can be attributed to variations in the unground samples. Like K, the amount of watersoluble P was largest in litter, intermediate in mor, and least in A1 horizon samples. Water-extracted P amounts found were similar for litter samples collected at the different locations (Fig. 1). However, the amount in the A1 horizon, and to a lesser extent from the mor, seemed related to the total quantity in the A1 horizon (Table 1) at that location.

The extracted water-soluble P content was largest from litter, mor, and A1 horizon samples heated to 200 C. For each kind of sample collected at locations 4, 5, and 6, the water-soluble P content decreased at higher temperatures. In contrast, A1 horizon samples from locations 1, 2, and 3 released similar amounts of water-soluble P following heating either above or below 200 C. Apparently, soilmineral differences at the 6 locations affected the release of water-soluble P from the samples that were heated to different temperatures.

The average percent of P in the water extracts, that is in the PO₄-ion form (rather than complexed with organic or inorganic soluble ions), is least for the litter and mor samples heated to 200 C than to higher or lower temperatures (Table 2). Presumably by heating to 200 C, nondestructive volatilization (Hosking, 1938) and structural changes preceding the 210 C endothermic reaction (Philpot, 1970) reduce the PO₄-ion solubility. In contrast, the A1 horizon water-soluble P had the smallest amount of uncomplexed P in the samples heated to less than 200 C and the greatest in samples heated to 200 C or higher. Dehydration may have destroyed the water-soluble complexingagent bonds to free the PO_{Δ} -ions.

The results of this laboratory study

probably can be used to predict the effect controlled burning has on nutrient release. Samples that were heated, either in the laboratory or in the controlled burn area, had similar physical characteristics (described previously for the laboratory study) if the temperatures were comparable. In the controlled burn area, the litter and mor were, respectively, 3 and 1 cm thick. Maximum temperatures were 343 C a few centimeters above the litter, 245 C in the litter, and 177 C in the mor. Mineral-soil maximum temperatures decreased from 165 to 65 C with depth. Thus, the detrimental effects of this controlled burn on nutrient release apparently would be negligible, except possibly in the litter if the temperature in the field and the laboratory can be compared directly.

Field-measured temperatures are an average temperature of a volume of soil or air, although the temperature of a plant fragment in the volume may be much higher when it burns. The temperatures of the laboratory heat treatments at 200 C or less are not subject to this error because the samples did not burn. At the 300 C temperature in the laboratory, some plant fragments burned more completely than others. This burning releases heat which increases the fragment temperature to something greater than the furnace temperature. The fragments in laboratory samples heated to 400 and 500 C had a more uniform color and fabric and presumably were heated to the same temperature. Although direct evidence is lacking, the physical appearance of the ash, partially burned plant fragments, and soil surface may be useful in assessing the effects of fire on the soil.

Prescribed burning as simulated in the laboratory may increase the quantity of water-soluble P and K in litter, mor, and upper A1 horizon of the soil if the temperature does not exceed 200 C. The total N content does not change until temperature exceeds 200 C. At higher temperatures, the water-soluble P and the total N content generally decreases for the litter, mor, and A1 horizon. The water-soluble K content in litter and mor decreased, but for most A1 horizons it

increased if heated to temperature greater than 200 C.

The appearance of the burned residue probably can be used as a guide to the effect of fire on total N content and the water-soluble P and K. Temperatures of 200 C darken and slightly char the organic fragments in the samples. From 200 to 400 C, the grayish-ash skeletal remains of plant fibers become more numerous, and between 400 and 500 C the combustible material is essentially destroyed, leaving a grayish ash residue for the organic-rich litter and mor and a reddish or grayish residue for the upper mineral A1 horizon.

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